

INTRODUCTION

Most of the Victoria quadrangle lies within an area that appears bright on telescopic images of the planet, the bright-albedo feature Aurora, which approximately coincides with the east half of the quadrangle (Davies and others, 1978, fig. 11). As is common with most of the imaged portions of Mercury, the Victoria quadrangle is dominated by basins and large craters, with plains materials occupying the areas between them. Almost all the pictures acquired by Mariner 10 that were used for mapping were obtained during the first encounter: those covering the southeast half of the quadrangle are incoming close-encounter images, and those covering the northwest corner are outgoing close-encounter images. At the time the pictures were obtained, the terminator was at about long 7° to 8°, within the eastern part of the quadrangle. A large gap in coverage between the incoming and outgoing images appears as a northeast-trending diagonal blank strip on the base map. A small part of this gap was filled in the southwestern part of the quadrangle by very poor second-encounter images.

No images provide a vertical view; in fact, the smallest angle between the planetary surface normal and the camera axis is about 50°. The high obliquity of the images, the wide range in sun-elevation angles, and the complete transection of the quadrangle by the gap in coverage greatly hamper geologic mapping. Only in about 15 percent of the quadrangle, near the southeast corner, do data permit separation of units with the confidence possible in other quadrangles on

STRATIGRAPHY

Three widespread units are recognized within the Victoria quadrangle. These are, from oldest to youngest, intercrater plains material, intermediate plains material, and smooth plains material. In addition, central peak, floor, rim, and ejecta materials related to the numerous craters and basins larger than about 20 km in diameter are mapped. The simplicity of the stratigraphic scheme is at least in part due to deficiencies in the data base; the history of plains formation almost certainly is more complex than our threefold division indicates, but we were not able to define consistent criteria of albedo, texture, and cratering for more than three plains units because of the highly variable quality of available pictures. PLAINS MATERIALS

About half of the intercrater area consists of material characterized by a very high density of small, mostly degraded craters, and an irregular to rough surface. Superposition relations suggest that this unit is about the same age as, or older than, all mappable craters and basins. The origin of intercrater plains material is enigmatic; some may be primitive crust, as implied by Trask and Guest (1975), but more likely it is of mixed origin, dominated by breccias formed by nowunrecognizable ancient craters. Some of the more plainslike areas included within this unit may well have an origin similar to that of intermediate plains material.

Within the 5° overlap area with the Kuiper quadrangle to the south, an area has been mapped that displays moderately rough to rough terrain and a high density of mostly degraded craters. This unit is very similar to intercrater plains material, and cannot be distinguished from it anywhere else in the Victoria quadrangle. Most of the cratered plains material is probably volcanic in origin, but some of it may consist of impact breccias.

Intermediate Plains Material Smooth to moderately irregular plains occupy most of the area between large craters not underlain by intercrater plains material. These plains superficially resemble the plains of the lunar maria; they generally have a relatively low albedo (Hapke and others, 1975) and are characterized by numerous elongate ridges. Like the lunar maria, the two younger mercurian plains units have been ascribed to volcanic activity (Trask and Guest, 1975; Strom and others, 1975; Trask and Strom, 1976), although this interpretation has been questioned (Wilhelms, 1976). A volcanic origin seems most probable, but no compelling evidence exists in the Victoria quadrangle to support this opinion.

The elongate ridges, though clearly associated with intermediate plains material, are not restricted to it. Locally, ridges extend into intercrater plains material adjacent to intermediate plains material, and large young (c4 and c3) craters superposed on the intermediate plains material commonly are transected by these Smooth Plains Material

Partly filling most craters is plains material that is smoother and less densely cratered than intermediate plains material. Because most areas underlain by this unit are enclosed within craters, contacts between smooth plains and older plains units are rare. Smooth plains material thus is defined almost entirely by texture and apparent crater density. Few superposition data directly support the inferred age sequence, but the relative youth of the smooth plains unit is indicated by its presence on the floors of craters that are superposed on intermediate plains material. The smooth plains unit probably includes materials of a wide range in age, but the exposed areas are too small to test this possibility quantitatively. Although a volcanic origin cannot be ruled out for all or part of the smooth plains material, it is more probably a mixture of ejecta from very small craters and colluvium mass wasted from crater walls.

STRUCTURE

The ridges associated with the intermediate plains unit are best interpreted as ectonic in origin because they extend into adjacent exposures of int plains material and, more significantly, because they transect ejecta, rims, and floors of craters. The ridges range in length from about 50 km to many hundreds of kilometers, are sinuous to lobate in plan, and generally trend about northsouth. Most are asymmetric, with one slope steeper than the other, and at places they can be more logically referred to as rounded scarps. Commonly, an individual ridge changes along trend from symmetric ridge to asymmetric ridge to rounded scarp. Strom and others (1975) interpreted most of these features to be surface expressions of thrust faults, and we can find no evidence within the Victoria quadrangle not already considered in their discussion. Because of their globally systematic orientations, these ridges and scarps have been associated with stresses developed by tidal despinning of Mercury (Melosh,

1977). However, most trend approximately north-south and thus do not fit the pattern expected in the midlatitude belt, unless stresses from overall contraction were superposed on the stresses due to despinning (Melosh, 1977, figs. 3 and 5). GEOLOGIC HISTORY The oldest material and features in the Victoria quadrangle are the intercrater plains material and areally associated, severely degraded basins. No craters are clearly older than intercrater plains material, and the relative ages of the c1

basins are ambiguous. Numerous large craters are superposed on intercrater plains material; by analogy with lunar and martian history (Hartmann, 1966, 1973; Neukum and others, 1975), these craters most likely date from more than about The available evidence suggests a relatively long history of plains formation. Some of the material included in the intercrater plains unit appears to have been plainslike before the intense cratering characteristic of the unit. In addition, the younger plains units exhibit densities of superposed craters ranging from moderate to very sparse. The intermediate plains material is older than the freshest (c4 and some c₃) large craters (100-150 km in diameter) but younger than all basins, and younger than all large craters that are more than moderately degraded (some c3, and all c2 and c1). Thus, the material mapped as the intermediate plains unit overlaps in time of origin the tail end of the primordial bombardment. The stresses responsible for the elongate ridges and scarps must have occurred after the end of the primordial bombardment and after emplacement of the intermediate plains unit. Where smooth plains material abuts ridges and scarps, the evidence is mostly ambiguous because we cannot tell if ridge formation involved smooth plains material or if the ridges are upwarped intermediate plains material with smooth plains material ponded against them. On the floors of some craters, such as Gluck, scarps apparently offset material mapped as smooth plains, but the exposures are so small that this interpretation could easily be challenged. Ridges appear to be both older and younger than medium-size craters (30-60 km in diameter) on the intermediate plains unit, but intersections of ridges with craters in this size range are too rare to constrain the time of ridge formation. Thus, ridge formation obviously occurred after emplacement of the intermediate plains unit, but how long after remains uncertain in this quadrangle. Smooth plains material is apparently younger than all large craters, and hence is the youngest material in the quadrangle with the exception of the local material

related to some very small craters (< 20 km in diameter). Figures 1 to 3 are plots of cumulative diameter versus frequency of craters on the three plains units. Figures 1 and 2 include craters \geq 3 km in diameter on the large areas of intercrater and intermediate plains units in the southeastern portion of the quadrangle. The two counting areas together are approximately bounded by lats 20° and 32° and longs 15° and 42° and are separated by an irregular but approximately east-west-trending contact; the intercrater plains material lies to

the south and the intermediate plains material to the north. Figure 3 includes craters > 1.2 km in diameter on smooth plains material that covers the floors of Because of the varied and generally poor quality of the imagery, detailed cratering history cannot be inferred from these plots. However, three observations seem valid: (1) the density of large craters is distinctly higher for intercrater than for intermediate plains material; (2) for the intercrater and intermediate plains material, curves for craters with diameters between 3 and 15 km nearly coincide (the abundant, mostly degraded small craters characteristic of intercrater plains material but not characteristic of intermediate plains material are less than 3 km in diameter); and (3) craters of all sizes on the smooth plains unit are much less abundant than on the other units, although the smooth-plains plot is unreliable in detail because of the small total number of craters counted and the need to combine counts from isolated exposures. A serious sampling problem exists for counting craters on the intermediate plains unit, because those in the diameter range of 50 to 150 km commonly occur in clusters, and it is very difficult to determine which craters of a cluster are younger than the surrounding plains unit and which are older. The area counted for figure 2 does not include any of these clusters. Immediately to the north, however, intermediate plains material surrounds three clusters of large craters, including Holbein, the c₄ crater ~ 100 km in diameter centered at lat 36° N., long 29°. This large crater clearly is younger than the surrounding intermediate plains material, and superposition relations suggest that two or three of the 50- to 60-km-diameter craters east of it also may be younger. The ages of other craters in the clusters, relative to the age of intermediate plains, are ambiguous. A second count of intermediate plains material was made that included the area of clusters of large craters, and that was based on the extreme assumption that all these large craters are younger than intermediate plains material. The resulting plot (not shown) differs from figure 2 at the large-crater end, as would be expected, but it still shows a lower density of large craters than does the plot of craters from intercrater plains material. Because not all the large craters occurring in scattered clusters are likely to be younger than intermediate plains material, the true plot of cumulative diameter versus frequency should not differ much from that shown

REFERENCES Davies, M. E., Dwornik, S. E., Gault, D. E., and Strom, R. G., 1978, Atlas of Mercury: National Aeronautics and Space Administration, Special Publication De Hon, R. A., Scott, D. H., and Underwood, J. R., Jr., 1981, Geologic map of the Kuiper quadrangle of Mercury: U.S. Geological Survey Miscellaneous nvestigations Series Map I-1233, scale 1:5,000,000. Hapke, B. W., Danielson, G. E., Jr., Klaasen, K. P., and Wilson, Lionel, 1975, Photometric observations of Mercury from Mariner 10, 1975: Journal of Geophysical Research, v. 80, no. 17, p. 2431-2443. Icarus, v. 5, no. 4, p. 406-418. ___1973, Martian cratering, 4, Mariner 9 initial analysis of cratering chronology Journal of Geophysical Research, v. 78, no. 20, p. 4096-4116. McCauley, J. F., Guest, J. E., Schaber, G. G., Trask, N. J., and Greeley, Ronald, 1981, Stratigraphy of the Caloris Basin, Mercury: Icarus, v. 47, no. 2, p. 184-202. Melosh, H. J., 1977, Global tectonics of a despun planet: Icarus, v. 31, no. 2, p. 221-243. Neukum, Gerhard, Konig, Beate, Fechtig, H., and Storzer, D., 1975, Cratering in

the Earth-Moon system: Consequences for age determination by crater counting: Lunar Science Conference, 6th Proceedings, p. 2597-2620. Strom, R. G., Trask, N. J., and Guest, J. E., 1975, Tectonism and volcanism on Mercury: Journal of Geophysical Research, v. 80, no. 17, p. 2478-2507. Trask, N. J., and Guest, J. E., 1975, Preliminary geologic terrain map of Mercury: Journal of Geophysical Research, v. 80, no. 17, p. 2461-2477. Trask, N. J., and Strom R. G., 1976, Additional evidence of mercurian volcanism: Icarus, v. 28, no. 4, p. 559-563. Wilhelms, D. E., 1976, Mercurian volcanism questioned: Icarus, v. 28, no. 4,

For sale by Branch of Distribution, U.S. Geological Survey, 1200 South Eads Street, Arlington, VA 22202, and Branch of Distribution U.S. Geological Survey, Box 25286, Federal Center, Denver, CO 80225